

PALEOBOTANY OF THE BIOGENIC UNIT OF THE LEFFE FORMATION (EARLY PLEISTOCENE, NORTHERN ITALY): BRIEF REPORT ON THE STATE- OF-THE-ART*

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RIASSUNTO - *Paleobotanica dell'unità biogenica della Formazione di Leffe (Pleistocene inferiore, Italia settentrionale): breve resoconto sullo stato dell'arte* - Il Quaternario *Italian Journal of Quaternary Sciences*, 8(2), 1995, 435-442 - Questo lavoro è un breve sommario dello stato dell'arte sulle ricerche condotte per lo studio del Bacino di Leffe (Italia settentrionale), dopo la costituzione, a Milano, di un apposito gruppo di lavoro per la revisione della serie di Leffe. E' stato eseguito un nuovo *record* pollinico, integrato con analisi preliminari sui macrofossili, delle ligniti e dei sedimenti lacustri che sono contenuti nella serie mediana di Leffe (la cosiddetta "unità biogenica" del Pleistocene inferiore). Vengono sottolineati i problemi che i dati paleobotanici pongono e che riguardano il paleoambiente, i tipi di migrazione della vegetazione, il paleoclima e la biostratigrafia degli inizi del Quaternario. Vengono citati i lavori specifici che trattano questi problemi in modo dettagliato. E' da notare come varie piante boschive, che risultano estinte nell'Europa centrale sin dal tardo Pliocene, erano ancora viventi nelle paludi di Leffe nei periodi più caldi degli inizi del Quaternario. La loro presenza risulta discontinua, modulata da cambiamenti climatici ciclici. I tradizionali criteri biostratigrafici, che sono basati sull'ipotesi di una progressiva estinzione delle piante terziarie, sono di difficile applicabilità quando si ha a che vedere con una tale evoluzione rapida e ciclica della vegetazione.

ABSTRACT - *Paleobotany of the biogenic unit of the Leffe Formation (early Pleistocene, N Italy): brief report on the state-of-the-art* - Il Quaternario *Italian Journal of Quaternary Sciences*, 8(2), 1995, 435-442 - This paper briefly summarises the state-of-the-art of the palaeobotanical research in the Leffe Basin (N-Italy), after constitution of a working group in Milan for the revision of the Leffe sequence. A new pollen record, integrated by preliminary macrofossil analyses, has been carried out in browncoal and lacustrine sediments in the middle Leffe sequence (the "biogenic unit" dating to the Early Pleistocene). Several problems arising from the palaeobotanical data are pointed out concerning the paleoenvironment, migration patterns of vegetation, palaeoclimatology and biostratigraphy at the beginning of the Quaternary. Reference is made to specific papers where these problems are examined in more detail. It is noteworthy that several woody plants, extinct in Central Europe since late Pliocene times, were still living in the Leffe mire during the warmest periods at the beginning of the Quaternary. Their record is discontinuous and it is modulated by cyclical climatic changes. Traditional biostratigraphic criteria, mainly based on a supposed progressive extinction of Tertiary plants, meet with substantial difficulties when applied to such rapid and cyclical evolution of vegetation.

Key words: Paleobotany, climatic changes, early Pleistocene, Leffe Formation, N Italy

Parole chiave: Paleobotanica, variazioni climatiche, Pleistocene inferiore, Formazione di Leffe, Italia settentrionale

1. INTRODUCTION

This paper summarises the first results of a biostratigraphic re-evaluation of the lacustrine and palustrine succession of the "biogenic unit" in the Leffe Basin sequence (Lombard Pre-Alps, N-Italy), based on new

palaeobotanical investigations⁽¹⁾ supported by a new geological survey and geomagnetic studies.

The geological evolution of the area during the Late Tertiary and Quaternary has been discussed in detail by Ravazzi (1993) and Cremaschi & Ravazzi (1995). Only a brief outline is presented here.

The Leffe Basin is located in a small valley, cut through Triassic carbonates and surrounded by 1,000-1,600 m high mountains, 15 km from the southern margin of the Pre-Alps (see Cremaschi & Ravazzi, 1995, fig. 1). The basin dates to Upper Pliocene times, during which the Po plain area was taken up by the sea, lapping the southern margin of the Lombard Pre-Alps.

The entire Leffe sequence consists of two complex stratigraphic units: the lower one (Leffe Formation) is composed of fan-delta, lacustrine and palustrine deposits, of Pleistocene age. The Leffe Formation is conformably overlaid by the *Gandino* Formation, composed of conglomerates and turbiditic lacustrine deposits dating to the Early Pleistocene.

(*) Paper presented to the Meeting "The significance of the Villafranchian in the Plio-Pleistocene stratigraphy" (Peveagno-Villafranca d'Asti, 24-29 June, 1994).

Lavoro presentato al Convegno "Il significato del Villafranchiano nella stratigrafia del Plio-Pleistocene", Peveagno-Villafranca d'Asti 24-29 giugno 1994.

(1) The paleobotanical researches constitute part of the research project "The continental succession of the Leffe Basin", among the programs of the CNR - Centro Geodinamica Alpina e Quaternaria in Milan. The financial support by the CNR Center, as well as by Museo Civico di Scienze Naturali di Bergamo, and by Comunità Montana Val Seriana, made it possible.

2. STRATIGRAPHIC POSITION OF THE BIOGENIC UNIT

The middle subunit of the Lefte Formation (biogenic unit) is especially important from a biostratigraphic point of view. It consists of a lacustrine and palustrine, 80 m-thick succession, made up of biogenic carbonate deposits, alternating with three organic layers (browncoal and *gyttja*). The middle one, more than 8 m thick, is called the "Main" browncoal layer.

On the basis of new paleomagnetic investigations (Billard *et al.*, 1983; Ravazzi *et al.*, 1992), the biogenic unit belongs to the Matuyama epoch, starting shortly after the upper reversal of the Olduvai event and ending well before the base of the Jaramillo event.

A rich fauna is included in the browncoal seams of the biogenic unit which has been known since the last century (Sordelli, 1896; Sthelin, 1930). This unit belongs to the Upper Villafranchian Mammal Age (Azzaroli *et al.*, 1986), and most probably to the Tasso or the Farneta Faunal Unit (according to Masini *et al.*, 1994 and Ambrosetti *et al.*, 1995). Another faunal assemblage, which includes *Elephas meridionalis vestinus*, was discovered in the Fifties during clay exploitation in the upper unit of the Lefte Formation which overlies the biogenic unit (Vialli, 1956). The stratigraphic position of this assemblage, previously attributed to the Cromerian (Vialli, 1956; Lona & Follieri, 1957), has been re-assigned to the Upper Villafranchian (Azzaroli *et al.*, 1986).

3. PALEOBOTANICAL INVESTIGATIONS IN THE BIOGENIC UNIT: PROJECT STRUCTURE

The pollen record by Lona (1950) in the biogenic unit outlined several climatic cycles. The new palynostratigraphic investigation has so far addressed the detailed description of a particularly well expressed cycle (Ravazzi & Rossignol Strick, 1995) in the 49±33 m interval of a new 186 m long core, taken in 1991 (Ravazzi *et al.*, 1992).

Only a few macroremains can be studied from this interval given the small quantity of material available. Consequently, several sections cropping at different stratigraphic positions have been sampled and tentatively correlated with the core. The main sites under study both for macro- and microfossils are:

- a) The main browncoal layer (8 m thick) excavated in 1992 at *Villa Giuseppina* site, from where coniferous wood has been studied (Ravazzi & Van der Burgh, 1994).
- b) Deposits with coniferous macroremains, dominated by *Tsuga*, buried by a slide along the paleoslope at the border of the Lefte Basin (Ravazzi *et al.*, in prep.).
- c) Fluvialite deposits interfingering with the lower unit of the Lefte Formation, where a rich carpoflora has been recovered (Casnigo site: Martinetto & Ravazzi, in preparation).

4. TAPHONOMY

The taphonomical processes involved in pollen dispersal and sedimentation, the catchment area of the Lefte Basin, the topography of the paleolake and diagenetic processes have been discussed by Ravazzi (1993, pp. 193-204). Moreover, a comparison between pollen and macroflora is in progress (Ghiotto *et al.*, in prep.; Pini, in prep.) to provide more insight into the distinction of the local, extralocal and regional components of the flora.

During the accumulation of the biogenic unit, the lake was isolated from the main river (Cremaschi & Ravazzi, 1995). Therefore, the regional component of pollen flora from the lacustrine and palustrine deposits was mostly supplied by a local wind circulation and by local rivers, draining a small, isolated catchment area (39 km²) surrounding the lake and bounded by high mountains. There is geological evidence that Tertiary deposits were absent in the catchment basin during the accumulation of the biogenic unit. Moreover, no sedimentary reworking structures have been observed (Ravazzi, 1993). Accordingly, reworked microfossils have not been recognized in these deposits.

In these taphonomic conditions, a change in pollen assemblages reflects vegetation development in the catchment area. Thanks to the isolated position of the lake, the comparison of a single pollen curve with the general evolution of the dominant vegetation can suggest whether it was autochthonous to the lake surroundings, or at least within the Lombard pre-alpine region (*taxa* such as *Aesculus*, *Nyssa*, *Rhododendron ponticum* type).

The macrofossil record helped to confirm that some plants were growing in the local (palustrine), or extralocal (wet plain, bottom valley woodlands) environment (e.g., *Aesculus*, *Glyptostroboxylon*, *Chamecyparis*, *Phellodendron*) (Ravazzi, 1994; Ravazzi & van der Burgh, 1994). However, in several cases, the only evidence available are high pollen concentrations (and estimated pollen influx), or a pollen record with a strictly local character (*i.e.*, hygrophilous plants whose pollen occurs exclusively in the browncoal). The case of *Cedrus* is a good example of the first situation; high values of pollen concentration (reaching 13,000 grains/g and an estimated pollen influx of 260 grains/cm²yr in marls of the biogenic unit) indicate that an extensive forest populated the mountains of the catchment area during sedimentation. However, macroremains have never been found in Lefte and only once in Italian Pleistocene deposits (Mastrorilli, 1965). *Cedrus* cone scales break away from the rachis near maturity; a single scale is very rarely transported from the dry habitat of this tree to a sedimentary basin (Spicer & Wolfe, 1987).

Evaluation of taphonomical processes as outlined above allowed the reconstruction of the local and regional vegetation (sections 5 and 6), and to understand the stratigraphic record of "Tertiary plants" section 7).

5. THE LOCAL VEGETATION

The pollen record of the biogenic unit is dominated throughout by arboreal *taxa* as well as local herbs and ferns growing *in situ*. The pollen diagrams are discussed in Ravazzi (1993) and Ravazzi & Rossignol Strick (1995) and only the vegetation reconstruction is presented here.

The local vegetation is especially important in the browncoal seams, which accumulated during phases of fen (herbaceous) and swamp (with tree layer) vegetation. Different types of communities growing in the mire have been documented on the basis of pollen, wood and browncoal composition.

Herbaceous communities

– Terrestrial or telmatic hygrophilous communities accumulating peat, dominated by *Phragmites*, or Cyperaceae (dominant *taxa*: *Carex sect. Acutae*, *Carex sect. Rostratae*, *Scirpus*, *Hypericum*). In some browncoal levels, mosses or ferns are dominant both in the palynoflora and in macroremains (*Thelypteris*, *Osmunda regalis*-type).

– Aquatic hydrophilous communities (*Potamogeton*, *Nymphaea*, *Nuphar*, *Sparganium minimum*). Pollen, seeds and trichosclereids of Nymphaeaceae (Pals *et al.*, 1980) are relatively abundant in peat and in *gyttja* deposits as well.

Swamps

– An eutrophic type with *Glyptostroboxylon tenerum* (Kraus) Conwentz, *Chamaecyparis*, *Nyssa*, *Pterocarya*, *Alnus* has been identified. This vegetation has been described from several Neogene browncoal sites in Central Europe [the Rhenish Basin (Teichmüller, 1958; Van der Burgh, 1987; Huhn *et al.*, 1994), the German-Polish depression (Schneider, 1990), the Cheb Basin in the Czech Republic (Buzek *et al.*, 1985)] and western North Italy (Martinetto, 1995). No Quaternary record in Europe has to date been identified.

– A vegetation type dominated by different species of *Picea*. *Picea* needles, together with Coniferous periderms and wood, has been found in distinct browncoal levels, and coincides with peaks of *Picea* pollen. From the pollen record, it appears that at least one species belonging to *P. excelsa*-type and another to the *P. omorika*-type appear to have been present in the local swamp vegetation. Macrofossils recovered in browncoal have been identified as *P. seriana* Sordelli (= *P. omorikoides* Weber ?), *P. balsami* Sordelli, *Picea cf. excelsa* foss. (Sordelli, 1896). A Late Pliocene swamp vegetation with *P. omorikoides* Weber has been described in Buzek *et al.* (1985) in the Czech Republic.

– An oligotrophic type vegetation, with *Sphagnum* spores and pollen peaks of *Pinus sylvestris*-type. Wood identified as *Pinus cf. tabulaeformis* Carrière has been

found in the same levels (Ravazzi & Van der Burgh, 1994).

The stratigraphic occurrence of these communities shows clear relationships with the development of the local vegetation, the regional vegetation, and climate (Ravazzi, 1993). For instance, mosses and ferns are never associated in the same stratigraphic interval, as are the first two types of swamp vegetation. Therefore, a floral list not-stratigraphically controlled from the Lefte browncoal (like that by Gregor, 1990) does not allow any deductions on the ecology of vegetation and paleoclimate to be made.

6. THE EVOLUTION OF THE REGIONAL VEGETATION AND CLIMATE CHANGE

The pollen record of extralocal and regional vegetation and the climate evolution is described in Ravazzi & Rossignol Strick (1995). It shows a repetitive forest succession of a mixed oak forest in a dry-temperate climate, followed by a Juglandaceae assemblage rich in Arcto-Tertiary elements, some of them indicating a wet climate (*Carya sp. pl.*, *Pterocarya*, *Fagus*, *Aesculus aff. hippostanum*), and a long phase of coniferous forests of cooler climate (*Picea*, *Tsuga*, *Cedrus*, *Abies*). A short phase of open vegetation (*Artemisia-Betula-Larix* pollen zones) interrupts this forest succession. It indicates a short event of drier and colder, continental climate.

By correlating the new pollen record with the data by Lona (1950), the complete vegetation dynamics seems to have been repeated at least seven times in the entire time interval represented by the biogenic unit.

Contrary to what outlined by Venzo (1950) and Lona (1950), the new pollen record and stratigraphic data show that these vegetation cycles cannot be related to the main glacial periods in the Alps described by Penck & Brückner (1909), but rather to high-frequency climatic cycles of moderate amplitude (Cremaschi & Ravazzi, 1995; Ravazzi & Rossignol Strick, 1995). A similar cyclicity has been observed in deep-sea cores of Early Pleistocene age (Ruddiman *et al.*, 1986; 1988; 1989) and in Mediterranean sapropel stratigraphy (Hilgen, 1987, 1991; Combourieu-Nebout, 1993; Lourens *et al.*, in press). Although the duration of the observed cyclicity in Lefte is still under study (Ravazzi & Moscardiello, submitted), a period of about 30 ka years can be estimated from counting laminated couplets in marls and from a calculation of sedimentation rate in peat after compression. This value is comparable to the period of the obliquity-driven cycles of 40 Ka.

7. THE CYCLIC RECORD OF RELICT PLANTS FROM TERTIARY

The new floral record from the biogenic unit points to the Quaternary persistence in the Southern Pre-Alps of several plants traditionally regarded as extinct in Eu-

rope at the end of the Tertiary, or surviving as relicts during the Early Pleistocene (Van der Hammen *et al.*, 1971). Most of these plants are occurring only during the warm-temperate and wet intervals of the Lefte record.

The following list includes exotic and other *taxa* of special interest for biostratigraphy, found in the biogenic unit and clearly autochthonous to the catchment area of the Lefte Basin. The criteria used to define autochthony have been briefly discussed in section 4 and are in agreement with criteria defined by the AIQUA Working Group on Quaternary Stratigraphy (AIQUA, 1995). Part of paleobotanical material is described in Ravazzi (1993; 1994) and in Ravazzi & Van der Burgh (1994). However, carpofloral identifications are still unpublished (Ghiotto *et al.*, in prep.).

The nomenclature follows Beug (1961), Moore *et al.* (1991); further taxonomical references are reported in the list.

Conifers

- Cedrus* aff. *atlantica* (pollen) (Montagnana, 1991; Ravazzi, 1993)
- Chamaecyparis* / *Thuja*-type (pollen)
- Chamaecyparis* (wood)
- Glyptostroboxylon tenerum* (Kraus) Conwentz (wood)
- Taxodium* type (pollen)
- Tsuga diversifolia* type (pollen)
- Tsuga* (cones and needles)

Angiosperms

- Acer palmatum* type (pollen) (Skawinska, 1985)
- Acer* Sect. *Saccharina* Pax (pollen) (Biesboer, 1975)
- Aesculus* aff. *hippocastanum* L. (pollen) (Ravazzi, 1994)
- Aesculus hippocastanum* (seeds) (Sordelli, 1896)
- Carya* spp. (pollen, fruits and wood)
- Celastraceae
- Eucommia* (pollen)
- Jasminum* (pollen)
- Juglans bergomensis* Massal. (fruits) [(Balsamo Crivelli) Massalongo, 1852]
- Juglans* Sect. *Trachycaryon* (pollen) (Stone & Broome, 1975) (= *J. cinerea*-type)
- Liquidambar* cf. *orientalis* Mill. (pollen)
- Magnolia cor* Ludwig (seeds)
- Nyssa* aff. *sylvatica* L. (pollen) (Lieux, 1983)
- Parthenocissus* (pollen)
- Phellodendron* (seeds)
- Pterocarya* (pollen, wood)
- Rhamnus* aff. *alaternus* L. (pollen)
- Rhododendron ponticum*-type (pollen)
- Staphylea* (pollen)

Taxa to be excluded

Cathaya bergeri (Kirchheimer) Schneider. The finds reported by Gregor (1990) are erroneous, because these cones, which originate from the collections of the

Museo Civico di Scienze Naturali di Bergamo, belong to *Pinus peuce* and are registered as coming from the Pianico Basin (Middle-Late Pleistocene), and not from Lefte.

Keteleeria (Lona & Bertoldi, 1973, table 1, fig. 1). The distinction between *Keteleeria* and *Abies* pollen is not documented by diagnostic characters.

The pollen record of most of the above listed *taxa* is discontinuous and modulated by climatic change. A cyclic, repetitive appearance-disappearance of a number of thermophilous woody plants, some of them characteristic for the mixed mesophytic forest (Braun, 1950; Wolfe, 1971) including several relict plants from the Tertiary, has been observed in the new pollen record (Ravazzi & Rossignol Strick, 1995), i.e., *Aesculus* aff. *hippocastanum*, *Carya*, *Celtis*, *Eucommia*, *Juglans* Sect. *Trachycaryon*, *Parthenocissus*, *Pterocarya*, and *Vitis*. Among "Tertiary plants", a group including Taxodiaceae e.g. *Glyptostrobus*, *Nyssa*, *Magnolia*, Cupressaceae e.g., *Chamaecyparis*, shows a more time-restricted occurrence, decreasing towards the top of the biogenic unit of the Lefte Formation.

Although these plants show a time-restricted record and low quantitative values, their repetitive appearance and disappearance suggests a migration regime from short-distance refugia to the Pre-Alps during periods of favourable climate. A discussion of the relationships between migration induced by climate changes, paleogeography and geological evolution of the southern alpine fringe at the beginning of the Quaternary requires an overview which cannot be provided. So far the marine/continental transition and the oldest fluvial deposits in the Lombard plain have not been dated, so that any attempt to correlate between the deposits in the valley floor of the Lombard Pre-Alps and in the plain is still unambiguous.

On the other hand, evidence for migration processes make it hard to develop stratigraphic inferences simply from the floral composition of a deposit. A temporary disappearance, produced by vegetation dynamics, can easily be misidentified as a stratigraphic extinction. Traditional biostratigraphic criteria, mainly based on the progressive extinction of "Tertiary relicts" (Van der Hammen *et al.*, 1971), meet with substantial difficulties when applied to such a cyclical and rapid floral evolution.

The biostratigraphic signal provided by the extinction of exotic plants may be used only in the context of large-scale biostratigraphy. For this purpose, a careful selection of irreversible events in a definite region has to be operated. The absence of *taxa* in the Lefte sediments whose regional extinction has been related by Martinetto (1995) with the cold phase at about 2.6 Ma, could also be relevant. None of those *taxa* with subtropical affinities, reported by Bertoldi *et al.* (1994) and Martinetto (1995) in Early and Middle Pliocene of N Italy, has been recognized in the Lefte floral record.

8. PLIO-PLEISTOCENE STRATIGRAPHY BASED ON FOSSIL PLANTS: A PROBLEMATIC TASK

The classic paleobotanical subdivision of the Pliocene and the Early Quaternary in Central Europe (Mai & Walther, 1988; Gibbard *et al.*, 1991; Zagwijn, 1992) and in the Mediterranean (Suc & Zagwijn, 1983; Suc, 1984; Bertoldi *et al.*, 1994) is based on a large scale time-constraint evolution of the flora and/or vegetation.

It is now clear that, with the high frequency glacial-interglacial couplets at about 3-2.5 Ma (Shakleton, 1995), these large-scale floral assemblages or/and "floral complexes" (*Florenkomplex*), defined by Mai & Walther (1988) as "a group of local fossil floras reflecting a chronologically delimited and distinguishable stage of vegetation", cannot describe biozones, but only a group of isolated points (the local fossil floras) in the stratigraphic record. Migration processes allow the same plants and vegetations to be in the same place before and after a "glacial" event, so that these points cannot be put in an unambiguous time sequence. Therefore, only few large-scale, yet irreversible events allow the establishment of biostratigraphic units as defined for local mammal faunas (De Giuli *et al.*, 1983), i.e. a group of local floras ranging after an irreversible event and before another. The record of local floras is discontinuous and its correlation potential is strongly limited in space. A synthesis of floral complexes, large-scale "climate-themes" (Zubakov & Borzenkova, 1990) and "vegetational provinces" has been attempted recently by Martinetto (1995) and Suc *et al.* (1995). The existence of "supercycles" during the whole Quaternary and even between 2.6 and 1 Ma has been recently asserted by Kukla (1995), based on loess sequences in China. However, this hypothesis has not yet been confirmed in deep sea records. Especially at the base of Early Pleistocene, there is no clear evidence of such large scale, irreversible events enabling to subdivide this time interval in a few stages. The climatic curve for the Pleistocene in The Netherlands (Zagwijn, 1992) shows that climatic oscillations, occurring at the boundaries separating stages (Tiglian, Eburonian, Waalian, Menapian, Bavelian, Cromerian) and during these stages, are of the same amplitude. Therefore, they have scarce potential of large scale correlation. Furthermore, the Leffe record suggests that a regime of high frequency (20÷40 ka) climatic cycles explain the floral evolution during the Early Pleistocene in Northern Italy. Stages longer than a cycle are not visible in the Leffe record so far investigated.

ACKNOWLEDGEMENTS

I wish to thank Dr. Edoardo Martinetto (Dip.to di Scienze della Terra, Università di Torino) for his contribution on macrofloral identification and comments on this text, Dr. Bas van Geel (Hugo de Vries Laboratory, University of Amsterdam), for introducing me on the analysis

of fungi and algal spores and Dr. Samuel P. Evans (Cranfield University) for improving the English form and for his useful comments.

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Ms. received: November, 1995

Final text received: December, 1995

Ms. ricevuto: Novembre 1995

Testo definitivo ricevuto: Dicembre 1995